

Manuscript prepared for Ocean Sci. Discuss.
with version 2014/07/29 7.12 Copernicus papers of the L^AT_EX class copernicus.cls.
Date: 20 March 2015

Modelling origin and transport fate of waste materials on the **south-eastern** southeastern Adriatic coast (Croatia)

M. Tudor¹ and I. Janeković²

¹Croatian Meteorological and Hydrological Service, Grič 3, Zagreb, Croatia

²Rudjer Bošković Institute, Bijenička 54, Zagreb, Croatia

Correspondence to: M. Tudor (tudor@cirus.dhz.hr)

Abstract

The ~~south-eastern~~ southeastern parts of the Adriatic Sea coastline were severely polluted by large amounts of accumulated waste material in the second half of November 2010. The waste, reported by major news agencies, accumulated dominantly during 21 November 2010 by favourable wind – ocean current transport system (the East Adriatic current is northwest so advects floating objects from southeast). In the study we analysed meteorological and oceanographic conditions that lead to the waste deposition using available in situ measurements, remote sensing data as well as numerical models of the ocean and the atmosphere. The measured data reveal that an intensive rainfall event from 7 till 10 November 2010, over the parts of Montenegro and Albania, was followed by a substantial increase of the river water levels indicating a possibility of flash floods that ~~possibly could have~~ splashed the waste material into a river and after to the Adriatic Sea ~~-(or to the sea directly)~~. The currents that can bring this waste to Croatia could have been intensified by the strong wind from southeast. In order to test ~~our hypothesis~~ these two hypotheses we set a number of numerical drifter experiments with trajectories initiated ~~off the coast of Albania over southeast Adriatic~~ during the intensive rainfall events following their faith in space and time. ~~One of the~~ The numerical drifter trajectory ~~experiment~~ experiments that resulted with drifters ~~reached that reached the~~ right position (~~south-eastern~~ southeastern Adriatic coast) and time (exactly by the time the waste was observed) ~~when were~~ initiated on 00:00 and 12:00 UTC of 10 November 2010 during the mentioned ~~flash flood~~ high precipitation event.

1 Introduction

On the 21 November 2010, a dramatic waste accumulation has been widely reported ~~at the south-eastern~~ by public media (web news agencies, television, radio, daily papers) at the southeastern coast of Croatia, particularly area of Pelješac Peninsula; islands Mljet, Korčula and Lastovo as well as in numerous inlets and beaches ~~in the vicinity~~ northwest of Dubrovnik (see map of the area in Fig. 1). The heaps of waste were composed mostly of

plastic packages, glass bottles, clothes and other typical floating municipal garbage while labels suggested that it's origin is Albania, Croatian neighbour country some part of the waste arrived from Albania. It is a country on the south-eastern southeastern Adriatic coast 100 km south-east southeast of the area where waste accumulated. The labels are no proof that the waste originated from Albania. One can easily imagine a heap of waste accumulated on a sea shore anywhere on Earth with labels "Made in China" since it is well known for its massive production and intensive export. This study describes a possible chain of events that lead to waste accumulation on beaches in southeast Croatia. It is not unusual that a few pieces of waste from Albania reach Croatian coast in a late autumn, however the event was several orders of magnitude larger than any other in previous years (according to the local officials, there were no reports in the media).

The meteorological conditions during October and November in southern Adriatic included several episodes of intensive precipitation that initiated flash floods in Montenegro and Croatia (there were no reports available for Albania). A flash flood event could have washed the waste to the sea (or first to a lake or a river that would eventually take it to the sea). There was no rainfall data from Albania available through standard international data exchange so remote sensing data and NWP model data were used to estimate which intensive precipitation events (if any) could have initiated a flash flood there. The high precipitation events that could have initiated flash flood in the area are identified by combination of in-situ and remote sensing data.

Numerical modeling studies that examine how a floating entity reached a certain position by means of atmosphere and sea driven currents have been done before (Beg Paklar et al., 2008; Döös et al., 2011; Liu and Weisberg, 2011). The subjects range from explanation of how floating sweet potato reached Polynesia from South America (Montenegro et al., 2008), spread of oil spills such as the one following the Deepwater Horizon disaster (see-) and has received more attention (for a collection of articles see <http://deepwaterhorizon.noaa.gov/>) as well as the floating debris that was washed to the sea by tsunami following the Tohoku 9 Mw earthquake earthquake on 11 March 2011 (see <http://www.marinedebris.noaa.gov/>).

In our case we use meteorological and ocean models to explore how waste items, once washed to the sea, floated to the area affected by the accumulated waste. During this study we assume that strong ~~south-easterly~~ southeasterly wind intensified the sea current system that was favourable to bring the floating waste materials ~~from Albania. Furthermore, one of the assumptions is~~. Previously, we also assumed that the waste was splashed into the sea by strong flash floods as a consequence of the severe torrential rain. ~~Those~~ These hypotheses are further investigated using all available meteorological, oceanographic and hydrological observational data as well advanced meteorological and oceanographic numerical models.

The next section describes the geographical characteristics ,of the area, meteorological and oceanographic conditions, measured data and models used in this study. ~~Results of~~ Results of the analysis of weather patterns using available measured data and results of model simulations are presented in Sect. 3 and summarized in ~~conclusions~~ conclusions in Sect. 4.

2 Region, data, and models

2.1 Region

The Adriatic Sea is a narrow sea, connected to the Mediterranean only by the Otranto Strait at the southern part. Bathymetry varies ~~over three orders of magnitude, with the northern part as~~ considerably over the basin. The northern part is the shallowest, with mean depth of 35 m; ~~the~~. The central and southern Adriatic are significantly deeper and divided by Palagruža sill ~~-(Fig. 1)~~. The central region reaches up to 280 m depth in Jabuka Pit. The southern region is the deepest, up to 1200 m in the South Adriatic Pit (SAP).

The Adriatic Sea surface flow is predominately cyclonic orientation (Cushman-Roisin et al., 2001) with distinct current regime of East Adriatic Current (EAC) flowing north-west along the eastern coast characterized with salty and warm water from the Ionian Sea. During the rain seasons EAC is further intensified with the outflow of the Albanian rivers

creating region of fresh water (ROFI) dynamics (Burrage et al., 2008, 2009). In the central region the sea surface flow typically bifurcates east of the Palagruza Sill (e.g. Wolf and Luksch, 1887) enhancing the cyclonic circulation in the southern Adriatic (Artegiani et al., 1997; Horton et al., 1997).

On the other side of the Adriatic Sea there is a Western Adriatic Current (WAC) holding fresher and colder water along the western coast. It carries a signature of Po river outflow, the most important source of fresh water in the whole Adriatic Sea.

On the land, the area is surrounded by Apennines in the west, Dinaric Alps and high mountains of Montenegro and Albania along eastern coast while on the northern coast reaches low and flat Po Valley. Mountains are much closer to the shore on the eastern side of the Adriatic Sea, with several peaks higher than 1.5 km located less than 10 km from the coast (Fig. 1). Those mountains have a strong effect on the air flow and atmospheric dynamics (Mesinger and Strickler, 1981) ~~defining and consequently define the~~ sea current response as well.

~~From the south,~~

Mediterranean cyclones often traverse the area (Horvath et al., 2009). However cyclones often form in the Genoa Bay, at the northwest (Mesinger and Strickler, 1981) traverse the Tyrhennian Sea and continue to the east possibly supporting cyclone development and intensification in the Adriatic Sea at the east and ~~Ionian-Ionian~~ Sea at the south (Alpert et al., 1990). These cyclones usually cross the Adriatic Sea but in a certain synoptic conditions can support development of a separate Adriatic cyclone (Horvath et al., 2008) and other mesoscale weather activity. In that case two of these cyclones can coexist forming a system of twin cyclones in which moist air converges and generates large quantities of available precipitable water (Lionello et al., 2006).

The intensive ~~dynamics found~~ atmospheric dynamics in the area also supports strong wind (Horvath et al., 2011; Bajić et al., 2007; Branković et al., 2008) development with the most severe and gusty wind from northeast named bura (see Grisogono and Belušić, 2009, for a review), as well the local wind from southeast ~~referred~~ referred as jugo (Jurčec et al., 1996). Strong bura or strong jugo can last for several days inducing strong response in

the Adriatic Sea (Kuzmić et al., 2006; Dorman et al., 2006). Onset, duration and spatial distribution of wind strength in a bura or jugo episode is controlled by an interaction of the synoptic and/or mesoscale forcing with local topography (Ivatek-Šahdan and Tudor, 2004; Pasarić et al., 2007; Tudor and Ivatek-Šahdan, 2010). The bura strength varies significantly in space and time forming usually several stripes of strong wind across the Adriatic Sea separated by areas of milder wind (Grubišić, 2004) related to mountain gaps and ridges upstream. On the other hand, jugo blows along shore, it is more steady and relatively warm wind related to a Genoa cyclone (Jurčec et al., 1996) or mesoscale cyclone above northern Adriatic (Brzović and Strelec Mahović, 1999; Brzović, 1999). Jugo can be associated with sirocco. However, the terms jugo and sirocco are not synonyms. The latter is a southern wind blowing from Sahara in advance of low pressure moving eastward across southern Mediterranean Sea. Jugo is SE wind over Adriatic, while in sirocco episodes, wind that brings the Sahara desert dust over Adriatic can be from south or southwest.

~~Climatologically~~Climatologically, Southern Adriatic region is characterized with warm and dry summers and mild and wet winters (Zaninović et al., 2008). The area receives abundant precipitation amounts as Crkvice in Montenegro holds the maximum measured on the European continent (Magaš, 2002). Parts of the northern Albania is rich in precipitation but unevenly distributed in space and time. In those regions precipitation can be further intensified by increased aerosol concentration (Koren et al., 2012) usually advected to the Adriatic Sea area from the Sahara desert by the sirocco wind.

It is worth to say that wind forcing, when pronounced, dominate over all other forcing contributions and dynamically shape the sea surface currents system found in the Adriatic Sea. The surface wind jets and wakes of the bura wind have a profound effect on the surface currents (Orlić et al., 1994; Pullen et al., 2003), while jugo wind is well known to influence WAC flow reversals (Orlić et al., 2007; Poulain et al., 2004). It is therefore important to drive the ocean model with a high resolution wind field that resolves high resolution wind features that develop due to interaction of large scale dynamics with local mountains surrounding the sea.

2.2 Data

In order to test our hypothesis and numerical model results we used available remote sensing data and in situ measurements. For the meteorological part we used SYNOP, climatological and rain-gauge measurements from Croatia, Montenegro, Italy, Greece and Macedonia. At the time of the event (November 2010) there were no in situ measured data available from Albania through standard meteorological network, the data from the airport did not contain measurements of precipitation. The hydrological analysis was based on the water level measurements on relevant major rivers in Montenegro and Macedonia used to confirm intensive ~~precipitation as possible flush flood events~~. precipitation as a possible cause of the flush flood event.

Remote sensing data, used in this study, originate from Meteosat Second Generation (MSG), specifically from The EUMETSAT Network of Satellite Application ~~Facilities~~ Facilities (NWC SAF)¹. Satellite derived precipitation data are used as provided from the Tropical Rainfall Measuring Mission (TRMM, Huffman et al., 2007), in particular we used the diurnal accumulated precipitation data from the 3B42 product and 3 hourly precipitation intensity data from 3B40RT product.

Precipitation can be enhanced by the presence of aerosols (Koren et al., 2012). The two sets of aerosol data presented in this study are the aerosol optical thickness (AOT) from Moderate Resolution Imaging Spectroradiometer (MODIS, Remer et al., 2008) aboard Aqua satellite and Ozone Monitoring Instrument aboard NASA's Earth Observing System (EOS) Aura satellite (OMI, Torres et al., 2002; Veihelmann et al., 2007)². The ~~meteorological model 10wind field is obtained by vertical interpolation from the lowest model level~~. ~~The model wind field quality was evaluated using the wind~~ wind over the sea surface derived from MetOp ASCAT (Bentamy et al., 2012; Bentamy and Croizé-Fillon, 2012) was used to evaluate 10 meter wind field from the meteorological model.

¹Products available on the <http://www.eumetrain.org/>.

²The OMI and TRMM data are available from Giovanni web server interface (Acker and Leptoukh, 2007) on <http://disc.sci.gsfc.nasa.gov>.

2.3 Models

2.3.1 Atmospheric model – ALADIN

The ~~NWP~~-numerical weather prediction (NWP) model data used in this study originate from the operational 8 km resolution forecast runs using ALADIN limited area model (Aire Limitée Adaptation Dynamique développement InterNational, ALADIN International Team, 1997) with a specific local 3-D-var data assimilation (Stanešić, 2011)~~that integrates~~. In autumn 2010, operational forecast run twice per day up to 72 h in advance starting from 00:00 and 12:00 UTC analyses. The model forecast in 8 km resolution ~~uses~~-used initial and boundary conditions from global model ARPEGE (Action de Recherche Petite Echelle Grande Echelle, Cassou and Terray, 2001) ~~used~~-run operationally in Meteo France. The operational high-resolution dynamical adaptation (Ivatek-Šahdan and Tudor, 2004) provides forecast of 10 m wind adapted to local and upstream topography (Horvath et al., 2011) in 2 km resolution. Unfortunately, this method provides only wind field at high resolution, but not the other meteorological variables needed to force the ocean model. The meteorological model 10 m wind field is obtained by vertical interpolation from the lowest model level (17 m above sea, see Geleyn, 1988, for more details).

In order to simulate the mesoscale characteristics and development of the low pressure field, a 2 km resolution forecast using the non-hydrostatic set-up of the ALADIN model and the full parametrization set, including radiation, microphysics and convection schemes (Tudor and Ivatek-Šahdan, 2010) was used to model the state of the atmosphere. The high-resolution forecast uses scale selective digital filter initialization Termonia (2008) and no data assimilation to initialize the model fields. It is coupled to the ALADIN 8 km resolution with 3 h interval. This might be insufficient to prevent the fastest of the meteorological features to enter the domain unnoticed by the lateral boundary coupling procedure (Tudor and Termonia, 2010) and possibly miss or undersample a storm rapidly entering the domain through the lateral boundaries. Since the southern Adriatic is not very far from the southern lateral boundary of the high resolution domain, model could have underestimated a storm

arriving from south through Otranto strait, however this would be a short duration event related to a flash flood but too short to affect the sea currents substantially.

2.3.2 Ocean model – ROMS

The ocean-quality of simulated currents on the ocean surface depends on the wind field. Wind field over Adriatic is variable in both space and time, depends on surrounding topography and events with strong and severe wind are better forecast in high resolution NWP models (Ivatek-Šahdan and Tudor, 2004; Branković et al., 2008; Tudor and Ivatek-Šahdan, 2010). The ocean dynamics as a response to the atmospheric forcing was computed using ROMS (Shchepetkin and McWilliams, 2005) Regional Ocean modelling System (ROMS, Shchepetkin and McWilliams (2005)) numerical model. ROMS model belongs to free surface, Boussinesq and hydrostatic approximation models that solves primitive equations using curvilinear finite difference grids. Model was forced with ALADIN meteorological model data (10 m wind, 2 m temp and relative humidity, sea level pressure, rainfall rate, short wave radiation and cloud fraction), climatological values for the Adriatic river run-offs and open boundary values with daily temperature, salinity, currents and sea level information from AREG (INGV) Mediterranean model. The advection scheme for tracers (temperature and salinity) is based on multidimensional positive definite advection transport algorithm – MPDATA (Smolarkiewicz and Margolin, 1998) while for momentum on 3rd order upwind scheme. More details of model implementation for the Adriatic Sea are described in Janeković et al. (2010). ROMS model time step was 120s while output frequency of needed current fields were stored every hour. Those hourly fields were used for computations of drifter trajectories.

The sea surface currents, responsible for waste transport, are computed using 2 km resolution ROMS ocean model and were used for virtual drifter trajectory simulations. Drifters are set to the surface layer, without vertical dynamics, ensuring representation of floating waste material. For computing numerical drifter trajectories we used Roms OFFline Floats (ROFF) package (Carr et al., 2008), available for ROMS community (http://web.atmos.ucla.edu/capet/Myresearch/my_research_floats.html).

3 Results and discussion

3.1 ~~Atmospherical~~ Atmospheric model results

To estimate the convective rainfall rate and precipitating ~~cloud~~ clouds we used derived fields from the NWC SAF products focused on studied area and time, and rain-gauge measurements and TRMM rainfall data. The NWC SAF precipitating clouds (PC, Thoss, 2012) field provides precipitation probabilities and the convective rainfall rate in mm/hour (CRR Rodriguez and Marcos, 2012) is computed assuming that clouds being both high and with a large vertical extent are more likely to induce rain (see <http://www.nwcsaf.org/> for more details). CRR gives estimate of intensive rainfall from convective clouds, but PC is useful estimate of rainfall from other types of clouds (eg. nimbostratus). According to the available rain-gauge measurements, 6 hourly PC and CRR fields and TRMM rainfall data, there were several heavy rainfall events in the month preceding 21 November 2010 that could have caused flash floods in the area of southeast Adriatic coast and inland.

We identify those events as four episodes: 23–25 October (E1), 2–3 (E2), 8–10 (E3) and 17–18 November 2010 (E4).

The large scale synoptic conditions responsible for meteorological setup are described using ERA Interim (Dee et al., 2011) re-analysis fields. It turns out that on 24 October 2010, low pressure system entered western Mediterranean from Atlantic, deepened and formed a cyclone, centred over Genoa bay. The next day the pressure decreased further and the associated southern wind strengthened from northern Africa to Adriatic causing intensive rainfall over the eastern Adriatic coast (E1). The cyclone moved southeast on 26 and initiated severe bura wind first on northern Adriatic and later spread over the whole Adriatic Sea by 27 October 2010. An ensemble of trajectories initiated ~~off the coast of Albania over~~ southeast Adriatic on 12:00 UTC, 25 October 2010 were used to test if this severe rainfall event was the one that flushed the waste to the sea ~~and consequently brought it to the~~ Croatian coast. The results of these trajectory computations are described later in the text (as experiment 2 in Sect. 3.3.1)

Another cyclone from the 1 to the 4 November 2010 (E2) moved from the Genoa bay southeastward, causing strong jugo wind over the Adriatic Sea (Fig. 6). The rainfall was the most intensive over the northern Italy and central Adriatic region with most of the rainfall above the Adriatic sea. Northern Adriatic received more than 100 mm of precipitation within 24 h, while the rain was weak in the southeastern region of our interest (Fig. 2). Consequently, E2 case was omitted from further analysis as was too weak to initiate a flash flood in southeast Adriatic. In the following days, meteorological situation was stable with weak pressure gradient, low wind as well high pressure over western Mediterranean inducing moderate winds from northwest.

The weather changed again in the period from 7 till 10 November 2010 (E3), dominated by a large scale cyclone (Fig 3) that arrived from the Northern Atlantic causing sirocco ~~over Mediterranean wind that brought~~ wind over Mediterranean (the colour of the wind vectors in Fig 3 indicates wind speed in m/s) that brought warm air and Sahara dust from the northern Africa (often found in rain gauges after intensive rainfall events), aerosols are also shown in Fig 3. Over Adriatic, the wind was strong to severe from southwest and south direction (Fig. 3). The wind direction was well forecast by the model, but at Palagruža, Dubrovnik, Prevlaka wind speed was underestimated, while at Mljet and Sv. Jure Biokovo the observed wind speed was correctly modelled. Pressure measurements reveal that during this event the Adriatic Sea was subject to a deep cyclone that last for several days (Fig. 6) with a strong pressure gradient over the Adriatic sea.

~~The-~~

~~The precipitation intensity was estimated using the~~ PC and CRR fields that showed strong convection and rainfall in the afternoon and evening ~~having periods with~~ with periods of weak to moderate rain intensity in the night and early morning. The precipitating clouds covered much of the area, while the convective rainfall rate is far more localized and very intensive.

It is important to note that PC and CRR fields were available on 6 h interval, while heavy rainfall could have occurred outside the sampling interval and easily could have been missed. The 24 hourly precipitation exceeds 100 mm over parts of southeast Croatia,

Bosnia and Herzegovina, Montenegro and northern Albania in TRMM precipitation estimates (shown as squares in Fig 4) and measurements at several rain-gauges in Montenegro for two consecutive days (rain-gauge measurements are shown as circles in Fig. 4) as well as in the model forecast (shaded background in Fig 4). Measurements from the rain-gauges showed that ~~rainfall during the during~~ E3, rainfall was the most intensive on stations in Montenegro (larger circles and stars on Fig. 2) hence on southeast Adriatic coast and significantly more intensive than in other episodes in November 2010. ~~This~~ For example, on 10 November 2010, there was 188.1 mm of rain measured at Cetinje and 143 mm measured at Golubovci station, both in Montenegro. Accumulated precipitation data are shown on maps for 9 and 10 November 2010 (for all available stations, circles in Fig 4). There were 3 intensive precipitation events before 21 November 2012, and measured precipitation exceeded 100 mm/24 hours in Montenegro only in the event 7 to 10 November 2012 (Fig 2. TRMM data also show that 24 hourly precipitation exceeded 100 mm over Albania in the same event. The forecast of the accumulated 24 hourly rainfall corresponds to the values measured on rain-gauges, although the model exaggerated slightly the rainfall on the coastline and underestimated the rainfall on several locations further inland (Fig. 4). Wind measurements (Fig. 6) show that wind in E3 episode was from south direction, more energetic and lasted longer than for other strong wind episodes during November 2010.

After E3, in the period from 11 till 15 November 2010, the weather was mostly dry with weak to moderate wind and direction typical ~~to~~ for the sea breeze diurnal cycle (Fig 6). In the next days a cyclone formed in the Genoa bay (15 November 2010) supporting again strong jugo wind over the whole Adriatic Sea (Fig. 5). Moreover, the wind strengthened (Fig. 6) with prevailing direction from southeast as measured in Dubrovnik, Mljet and Prevlaka (Fig. 7). Precipitation was intensive with peaks above 100 mm within 24 h on 17 November (E4), but the maxima were localized on northeast Adriatic ~~region, outside of our scope area~~. The ALADIN model forecast had similar rainfall distribution, as a result, E4 was omitted from detailed analysis as a period favourable with respect to the flash flood. However, during it is important as the wind field has driven the currents on the sea surface. During E4 wind was stronger on the coast (Dubrovnik and Prevlaka) than in the off-shore region (Mljet), as a con-

sequence of channelling effect the coastal mountains ~~-(Fig 7)~~. Model yields stronger wind above open sea (thinner arrows on regular grid in Fig 7) over southern Adriatic than MetOp ASCAT wind data (thicker arrows in Fig 7) for 16 November 2010, but the wind strength and direction are correct for 17 November 2010 (Fig 7). The global pressure gradient over the Mediterranean and Central Europe supported the wind regime from south and southeast over the whole southern Adriatic (Fig 5). Later, by 19 November 2010, the wind changed direction to southwest. A cyclone moved from Atlantic southeast, to the western Mediterranean. The wind changed to strong and severe jugo wind on the 21 November 2010.

Based ~~only on the previous analysis on the analysis of available precipitation fields~~, it turns out that E3 episode was the one when intensive rainfall occurred over ~~Albania and coast and inland of southeast Adriatic and~~ was the most likely event that could have triggered a flash flood there.

3.2 Hydrology

Annual river run-off distribution for the Albanian rivers usually varies for an order of magnitude during the year (embedded diagram in Fig. 8) with one pronounced peak in November and another in January. Bojana river collects the water flowing from Drim river and Skadar lake and flows into Adriatic along the border between Albania and Montenegro.

The largest lake in the region, the Skadar lake, is filled by river Morača and Crnojevica in Montenegro and drained into Bojana River (the name is Bojana in Montenegro and Buna in Albania). Bojana River also receives Drim river as a major tributary on the way to the Adriatic Sea. Drim (Drim in Montenegro, Drin in Albania) river powers 3 hydroelectric power plants in Albania. Downstream it splits into two flows, the smaller one reaches the Adriatic sea directly, and the larger part flows into Bojana River.

The river water level measurements on the rivers in Montenegro that belong to the Adriatic Sea catchment (Fig. 8) increase ~~substantially~~ substantially for the E3 episode. The water level surge is most intense for the rivers that fill the Skadar lake. The level of Bojana River raised as well during the same event. This is followed by a rise of 1.5 m in the water level of the Skadar lake. Bojana River level rose before the level of ~~Lage Skadar~~ Skadar Lake, this

could have happened due to an increase in contribution from the Drim river tributary. The water levels of Skadar lake and Bojana River stayed high until the end of November 2010.

Those measurements suggest that the event (E3) from 8 till 10 November 2010 was capable of flushing the waste material into the Adriatic sea or any of the rivers in the area that flow into it.

3.3 Ocean model results

Ocean model results show (Fig 9) consistent development of strong surface northwest currents after strong jugo wind episodes and small eddies close to the eastern Adriatic coast in the periods of weak wind forcing. As stated before in the text, during the November 2010, we can find three periods of different wind conditions over the Adriatic Sea. The first one from 7 till 11, when strong south-east southeast wind generated strong north-west current system-northwest current system in southeast Adriatic (Fig. 9a) followed by a weak wind period when the sea-current transport was weaker (Fig. 9b) and the ocean model formed a pool of colder water in southeast Adriatic in the area where Bojana river enters the Adriatic Sea. Finally, the period with moderate to strong south-east southeast wind (Fig. 9c) that strengthened the north-west strengthened the northwest current. This event was most likely responsible for waste transport and deposition.

Numerical drifter trajectories

~~In order to test hypothesis we set a~~The waste consisted of floating items, so its movement was computed as virtual floating drifters released in southeast Adriatic. Trajectories of virtual drifters were computed using ROFF package and surface currents from 2 km resolution ROMS run. The computation of the trajectories of the drifters stopped when they reached coastline for the first time. The drifters would then stay at this position on the coastline. This allowed accumulation of drifters.

A number of numerical drifter experiments was set in which trajectories were initiated of the coast of Albania in southeast Adriatic on the 12:00 UTC, 19 (experiment 1) and 25

October 2010 (experiment 2), and then sequentially at 00:00 and 12:00 UTC on each day starting from 8 till 12 November 2010 (experiments 3–11). All virtual drifters were released within a polygon covering an area ~~ver southern Adriatic close to Albanian coastline over southeastern Adriatic~~. The initial points of virtual drifter trajectories are separated by 0.01° (≈ 1 km) along longitude and latitude (total 3071 drifters) and fill a polygon with longitude and latitude coordinates of southwest corner (19.1,41.0) and northeast corner (19.4,41.9) that cover a portion of southeastern Adriatic Sea in the vicinity of coast of Albania and Montenegro (Fig 10). Furthermore we divided the polygon into 9 areas (A1, . . . , A9) to better cluster track different regions, hence possible source origin. The drifters starting from different areas are plotted in different colours, as marked on the Fig. 10. The longitude and latitude coordinates of southwest and northeast points of the polygons are

- A1: SW (19.1,41.0), NE (19.2,41.3) shown in red,
- A2: SW (19.2,41.0), NE (19.3,41.3) shown in green,
- A2: SW (19.2,41.0), NE (19.3,41.3) shown in green,
- A3: SW (19.3,41.0), NE (19.4,41.3) shown in blue,
- A4: SW (19.1,41.3), NE (19.2,41.6) shown in magenta,
- A5: SW (19.2,41.3), NE (19.3,41.6) shown in cyan,
- A6: SW (19.3,41.3), NE (19.4,41.6) shown in yellow,
- A7: SW (19.1,41.6), NE (19.2,41.9) shown in dark green,
- A8: SW (19.2,41.6), NE (19.3,41.9) shown in orange,
- A9: SW (19.3,41.6), NE (19.4,41.9) shown in gray,

A plot of drifter positions was done with 6 hourly interval for each experiment (not shown), and the summary trajectories are shown in Fig 10.

- In the ~~case of~~ experiment 1 the drifters were released at 12:00 UTC, 19 October 2010 and were first pushed offshore into EAC. It turns out that a considerable number of drifters originated from regions A7, A8 and A9 reached Croatian coast and Mljet island already on 27 October 2010. The ~~rest of~~ drifters from regions A4, A5 and A6 reached Mljet channel by 3 November, but were pushed back ~~south-east~~ southeast in the following days. Those drifters continued further to the ~~north-west~~ northwest and finally accumulated on the islands much further northwest than observed (Fig. 10a). There were no reports of significant accumulation of waste on the Croatian coast that would be a consequence of this event. In that sense, we can reject the hypothesis that this rainfall event was the one that caused ~~flash flood and~~ the flash flood that got the waste material to the sea.
- For experiment 2 the drifters were released at 12:00 UTC on 26 October 2010. Soon the ~~drifters~~ drifters were advected in the westward direction. When drifters entered ~~WAGEAC~~, they moved more to the ~~north-west~~ northwest and were deposited on Mljet Island already on 9 November 2010. However, several drifters starting from A6 region entered Mljet channel on 18 and later deposited on Pelješac on 21 November (Fig. 10b). Based on those results we can assume that it is possible but unlikely that the rainfall event on 26 October 2010 has initiated the chain of events that led to severe waste disposal in the region.
- Drifters initiated on 8 November (both 00:00 and 12:00 UTC – experiments 3 and 4) mostly arrived to ~~south-east~~ southeast Adriatic Sea coast, the northern Albania and Montenegro as soon as on the 11 November 2010 (Fig. 10c and d) as a consequence of sea current system supported by strong southern and SSW wind blowing on 8 and 9 November 2010. Furthermore, strong wind changed direction into NW on 12 and 13 November 2010 generating currents that transported numerical drifters off the coast, resulting only with a small number of them (initiated from A4 region), to reach Mljet island and coastline ~~north-west~~ northwest of Dubrovnik by 18 November 2010. The rest of the drifters dominantly stayed in the ~~south-east~~ southeast region, while

only a small number of them moved ~~north-eastward~~ northeastward not entering Mljet Channel, but instead floated much closer to coast, at the end finally accumulated in the Koločep Channel and Ston bay area on 25 November 2010.

- ~~Quickest~~ The quickest drifters initiated from A7, A4 and A8 regions at 00:00 and 12:00 UTC on 9 November 2010 (experiments 5 and 6) reached Mljet Island and entered Mljet Channel already on 17 and 18, while a majority of drifters from other areas accumulated in the Ston bay after 21 November 2010 (Fig. 10e and f).
- ~~Small~~ A small number of drifters from A7, A8 and A9 regions, released at 00:00 and 12:00 UTC on 10 November 2010 (experiments 7 and 8) reached Mljet, Dubrovnik and Koločep channel by 18, while other drifters initiated from the same area accumulate on the Croatian shores on 21 November 2010 (Fig. 10g and h). ~~In the case of~~ The drifters initiated further south lagged behind former ones and approached the affected area on 22 November 2010. Those dates were ~~the most commonly reported~~ reported in the media as the onset of severe pollution at the Croatian coast.
- The last three sets of drifters, released at 00:00 and 12:00 UTC on 11 and 00:00 UTC on 12 November 2010 (experiments 9–11) were first pushed westward off the Albanian coast and stayed in the area off shore of Albania and Montenegro for a few days. Later they were transported into southward direction on 14 to 16 November ~~2012~~ 2010 (Fig. 10i). Apparently, EAC was detached, at that time, from the shore and it's typical path. As a result, drifters from A1 and A2 regions arrived further ~~north-westward~~ northwestward than drifters initiated more to the north or closer to the coastline.

One should bear in mind that the drifter trajectories do not allow ~~as us~~ to assign a single event in space and time as the moment when the waste was disposed to the sea. There is ambiguity in the position as a result of unresolved physics, ~~unperfect meteorological model~~ imperfect meteorological model and initial conditions used to force the ocean counterpart, missing dynamics in the ocean model introduced with a lack of ~~wave-current~~ wavecurrent interaction, spatial model resolutions in the narrow channels etc. However, the results do

show that there is a possibility and is the most probably that the heavy rain on 9 and 10 November 2010 washed the waste into the sea (or first to a river that carried it to the sea by that date). The computations further show that not all the ~~waste that was washed into the sea from the Albanian shore~~ numerical drifters initiated in southeast Adriatic inevitably ended on the coastline of ~~south-east of southeast~~ Croatia. Surface sea currents enhanced by the wind forcing can carry the waste back to the shore, or to the closer coastline of Montenegro. Otherwise, different meteo-ocean conditions can push the waste off shore, and EAC can carry the waste to central or even north Adriatic, or in some cases ~~even into~~ back to the southern regions of the Adriatic Sea. However, none of the trajectories initiated in our experiments crossed the Adriatic Sea and approached to Italy, which is probably due to an absence of intensive bura events during the studied ~~periods~~period.

4 Conclusions

The oceanographic and meteorological conditions that lead to a severe deposition of waste material on the ~~south-eastern~~ southeastern Adriatic Sea coast on 21 November 2010, are studied using ALADIN – meteorological and ROMS – ocean numerical models along with ~~all~~ available measurements. ~~Given the fact that on labels from the part of the retrieved waste indicated an Albanian origin, we~~ We tried to answer what, where and when was the cause for the event. The initial points presented in this study were limited to the area of southeast Adriatic since the labels indicated that some of the items possibly originate from Albania and the local current system is northwest.

Based on the meteorological simulations and satellite derived precipitation reveal several intensive rainfall events that could have initiated flash floods in Albania and presumably flush the waste material to the rivers and later to the Adriatic Sea. Moreover, measured and NWP model rainfall data shows that the rain was more intensive over the Albania in the event from 8 till 10 November 2010 (E3) than in the other intensive rainfall events that occurred in the studied area during the 4 weeks before the reported waste accumulation.

Measured wind speed during the episode E3 was strong to severe from southern direction, however slightly underestimated by the operational ALADIN model forecast at several land locations. Improvements in the atmospheric model resolution could resolve those issues as noted in Signell et al. (2005). [Since strong wind influences the surface currents that advect the drifters, this could have an impact on the computed trajectories of virtual drifters.](#) It is interesting to note that based on the ASCAT estimated wind data, for the 16 November 2010, ALADIN model wind speed was larger than the measured one ([Fig 7](#)). During the last studied period (E4), wind from observations as well from the model, was from [south-east-southeast](#) with weaker magnitudes than in the E3 period. In the E4 case the strongest wind was found over the open sea, in the [south-east-southeast](#) region of Mljet Island as well south of Dubrovnik ([Fig 7](#)). This event is a typical jugo wind episode which further enhanced the sea surface current system – responsible for transport of the waste material to the Croatian waters and finally to the shore on 21 November 2010.

Acknowledgements. The authors would like to thank the Meteorological and Hydrological Service of Montenegro for providing the meteorological and water level measurements as well AREG – INGV, Bologna, Italy [center for boundary conditions of ROMS model](#). The authors are grateful to NASA for providing valuable satellite derived products through the GIOVANNI web interface as well as TRMM, OMI and MODIS scientists and developers. The NOAA NESDIS CoastWatch and NOAA SWFSC ERD are acknowledged for providing Metop ASCAT wind data. [Dr. I. Janeković](#) was sponsored through project [098-0982705-2707 of Croatian Ministry of Science](#) [HRZZ-5928 of Croatian Science Foundation](#). Martina Tudor thanks the Croatian Ministry of Science for support through the grant 004-1193086-3036.

References

- Acker, J. G. and Leptoukh, G.: Online analysis enhances use of NASA earth science data, *Eos Trans. AGU*, 88, 14–17, 2007.
- ALADIN International Team: The ALADIN project: Mesoscale modelling seen as a basic tool for weather forecasting and Atmos. Research, *WMO Bull.*, 46, 317–324, 1997.

- Alpert, P., Neeman, B. U., and Shay-El, Y.: Intermonthly variability of cyclone tracks in the Mediterranean, *J. Climate*, 3, 1474–1478, 1990.
- Artegiani, A., Bregant, D., Paschini, E., Pinardi, N., Raicich, F., and Russo, A.: The Adriatic Sea general circulation, Part II: Baroclinic circulation structure, *J. Phys. Oceanogr.*, 27, 1515–1532, 1997.
- Bajić, A., Ivatek-Šahdan, S., and Horvath, K.: Spatial distribution of wind speed in Croatia obtained using the ALADIN model, *Cro. Met. J.*, 42, 67–77, 2007.
- [Beg Paklar, G., Žagar, N., Žagar, M., Vellore, R., Koračin, D., Poulain, P.-M., Orlić, M., Vilibić, I., and Dadić, V.: Modeling the trajectories of satellite-tracked drifters in the Adriatic Sea during a summertime bora event, *J. Geophys. Res.*, 113, C11S04, doi:10.1029/2007JC004536, 2008.](#)
- Bentamy, A. and Croizé-Fillon, D.: Gridded surface wind fields from Metop/ASCAT measurements, *Int. J. Remote Sens.*, 33, 1729–1754, 2012.
- Bentamy, A., Grodsky, S. A., Carton, J. A., Croizé-Fillon, D., and Chapron, B.: Matching ASCAT and QuikSCAT winds, *J. Geophys. Res.* 117, C02011, doi:10.1029/2011JC007479, 2012.
- Branković, Č., Matjačić, B., Ivatek-Šahdan, S., and Buizza, R.: Downscaling of ECMWF ensemble forecasts for cases of severe weather: ensemble statistics and cluster analysis, *Mon. Weather Rev.*, 136, 3323–3342, 2008.
- Brzović, N.: Factors affecting the Adriatic cyclone and associated windstorms, *Contr. Atmos. Phys.*, 72, 51–65, 1999.
- Brzović, N. and Strelec Mahović, N.: Cyclonic activity and severe jugo in the Adriatic, *Phys. Chem. Earth*, 24, 653–657, 1999.
- Burrage, D., Wesson, J., Martinez, C., Pérez, C., Möller, O., and Piola, A.: Patos Lagoon outflow within the Río de la Plata plume using an airborne salinity mapper: observing an embedded plume, *Cont. Shelf Res.*, 28, 1625–1638, 2008.
- Burrage, D. M., Book, J. W., and Martin, P. J.: Eddies and filaments of the Western Adriatic Current near Cape Gargano: analysis and prediction, *J. Marine Syst.*, 78, 205–226, 2009.
- [Carr, S.D., Capet, X.J., McWilliams, J.C., Pennington, J.T., and Chavez, F.P.: The influence of diel vertical migration on zooplankton transport and recruitment in an upwelling region: Estimates from a coupled behavioral-physical model. *Fisheries Ocean.*, 17, 2008.](#)
- Cassou, C. and Terray, L.: Oceanic forcing of the wintertime low-frequency atmospheric variability in the north atlantic european sector: a study with the arpege model, *J. Climate*, 14, 4266–4291, doi:10.1175/1520-0442(2001)014<4266:OFOTWL>2.0.CO;2, 2001.

- Cushman-Roisin, B., Gačić, M., Poulain, P.-M., and Artegiani, A. (Eds.): *Physical Oceanography of the Adriatic Sea*, Kluwer Academic Publishers, Dordrecht., 2001.
- Dee, D. P., Uppala, S. M., Simmons, A. J., Berrisford, P., Poli, P., Kobayashi, S., Andrae, U., Balmaseda, M. A., Balsamo, G., Bauer, P., Bechtold, P., Beljaars, A. C. M., van de Berg, L., Bidlot, J., Bormann, N., Delsol, C., Dragani, R., Fuentes, M., Geer, A. J., Haimberger, L., Healy, S. B., Hersbach, H., Hólm, E. V., Isaksen, I., Kållberg, P., Köhler, M., Matricardi, M., McNally, A. P., Monge-Sanz, B. M., Morcrette, J.-J., Park, B.-K., Peubey, C., de Rosnay, P., Tavolato, C., Thépaut, J.-N., and Vitart, F.: The ERA-Interim reanalysis: configuration and performance of the data assimilation system, *Q. J. Roy. Meteor. Soc.*, 137, 553–597, doi:10.1002/qj.828, 2011.
- [Döös, K., Rupolo, V., Brodeau, L.: Dispersion of surface drifters and model-simulated trajectories, *Ocean Modelling*, 39, 3-4, 301-310, 2011.](#)
- Dorman, C., Book, J., Carniel, S., Cavaleri, L., Chiggiato, J., Doyle, J., Grbec, B., Haack, T., Janeković, I., Lee, C., Malačić, V., Orlić, M., Pullen, J., Russo, N., Paschini, E., Sclavo, M., Vilibić, I.: Winter 2003 marine atmospheric conditions and the Bora over the Northern Adriatic, *J. Geophys. Res.*, 111, C03S03, doi:10.1029/2005JC003134, 2006.
- Geleyn, J.-F.: Interpolation of wind, temperature and humidity values from model levels to the height of measurement, *Tellus A*, 40, 347–351, 1988.
- Grisogono, B. and Belušić, D.: A review of recent advances in understanding the meso- and microscale properties of the severe Bora wind, *Tellus A*, 61, 1–16, 2009.
- Grubišić, V.: Bora-driven potential vorticity banners over the Adriatic, *Q. J. Roy. Meteor. Soc.*, 130, 2571–2603, 2004.
- Horton, C., Clifford, M., Schmitz, J., and Kantha, L. H.: A real-time oceanographic nowcast/forecast system for the Mediterranean Sea, *J. Geophys. Res.*, 102, 25123–25156, 1997.
- Horvath, K., Lin, Y.-H., and Ivančan-Picek, B.: Classification of cyclone tracks over the Apennines and the Adriatic Sea, *Mon. Weather Rev.*, 136, 2210–2227, 2008.
- Horvath, K., Ivatek-Šahdan, S., Ivančan-Picek, B., and Grubišić, V.: Evolution and structure of two severe cyclonic bora events: contrast between the northern and southern Adriatic, *Weather Forecast.*, 24, 946–964, 2009.
- Horvath, K., Bajić, A., and Ivatek-Šahdan, S.: Dynamical downscaling of wind speed in complex terrain prone to bora-type flows, *J. Appl. Meteorol. Climatol.*, 50, 1676–1691, 2011.
- Huffman, G. J., Bolvin, D. T., Nelkin, E. J., Wolff, D. B., Adler, R. F., Gu, G., Hong, Y., Bowman, K. P., and Stocker, E. F.: The TRMM Multisatellite Precipitation Analysis (TMPA): quasi-global, multiyear, combined-sensor precipitation estimates at fine scales, *J. Hydrometeorol.*, 8, 38–55, 2007.

- Ivatek-Šahdan, S. and Tudor, M.: Use of high-resolution dynamical adaptation in operational suite and research impact studies, *Meteorol. Z.*, 13, 1–10, 2004.
- Janeković, I., Dutour Sikirić, M. A., Tomažić, I., and Kuzmić, M.: Hindcasting the Adriatic Sea surface temperature and salinity: a recent modeling experience, *Geofzika*, 27, 85–100, 2010.
- Jurčec, V., Ivančan-Picek, B., Tutiš, V., and Vukičević, V.: Severe Adriatic jugo wind, *Meteorol. Z.*, 5, 67–75, 1996.
- Koren, I., Altaratz, O., Remer, L. A., Feingold, G., Martins, J. V., and Heiblum, R. H.: Aerosol-induced intensification of rain from the tropics to the midlatitudes, *Nat. Geosci.*, 5, 118–122, 2012.
- Kuzmić, M., Janeković, I., Book, J. W., Martin, P. J., and Doyle, J. D.: Modeling the northern Adriatic double-gyre response to intense bora wind: a revisit, *J. Geophys. Res.-Oceans*, 111, C03S12, doi:10.1029/2005JC003377, 2006.
- Lionello, P., Bhend, J., Buzzi, A., Della-Marta, P. M., Krichak, S. O., Jansà, A., Maheras, P., Sanna, A., Trigo, I. F., and Trigo, R.: Cyclones in the Mediterranean region: climatology and effects on the environment, in: *Mediterranean Climate Variability*, edited by: Lionello, P., Malanotte-Rizzoli, P., and Boscolo, R., 325–372, Elsevier, 2006.
- [Liu, Y., and Weisberg R. H.: Evaluation of trajectory modeling in different dynamic regions using normalized cumulative Lagrangian separation, *J. Geophys. Res.*, 116, C09013, doi:10.1029/2010JC006837, 2011.](#)
- Magaš, D.: Natural-geographic characteristics of the Bok Kotorska area as the basis of development, *Geoadria*, 7/1, 51–81, 2002.
- Mesinger, F. and Strickler, R. F.: Effect of mountains on Genoa cyclogenesis, *J. Meteorol. Soc. Jpn.*, 60, 326–337, 1981.
- Montenegro, A., Avis, C., and Weaver, A. J.: Modelling the pre-historic arrival of the sweet potato in Polynesia, *J. Archaeol. Sci.*, 35, 355–367, 2008.
- Orlić, M., Kuzmić, M., and Pasarić, Z.: Response of the Adriatic Sea to the bora and sirocco forcing, *Cont. Shelf Res.*, 14, 91–116, 1994.
- Orlić, M., Dadić, V., Grbec, B., Leder, N., Marki, A., Matić, F., Mihanović, H., Beg Paklar, G., Pasarić, M., Pasarić, Z., and Vilibić, I.: Wintertime buoyancy forcing, changing seawater properties and two different circulation systems produced in the Adriatic, *J. Geophys. Res.-Oceans*, 112, C03S07, doi:10.1029/2005JC003271, 2007.
- Pasarić, Z., Belušić, D., and Bencetić Klaić, Z.: Orographic influences on the Adriatic sirocco wind, *Ann. Geophys.-Italy*, 25, 1263–1267, 2007.

- Poulain, P.-M., Mauri, E., and Ursella, L.: Unusual upwelling event and current reversal off the Italian Adriatic coast in summer 2003, *Geophys. Res. Lett.*, 31, L05303, doi:10.1029/2003GL019121, 2004.
- Pullen, J., Doyle, J. D., Hodur, R., Ogston, A., Book, J. W., Perkins, H., and Signell, R.: Coupled ocean–atmosphere nested modelling of the Adriatic Sea during winter and spring 2001, *J. Geophys. Res.*, 108, 3320, doi:10.1029/2003JC001780, 2003.
- Remer, L. A., Kleidman, R. G., Levy, R. C., Kaufman, Y. J., Tanré, D., Matoo, S., Vanderlei Martins, J., Ichoku, C., Koren, I., Yu, H., and Holben, B. N.: Global aerosol climatology from MODIS satellite sensors, *J. Geophys. Res.*, 113, D14S07, doi:10.1029/2007JD009661, 2008.
- Rodriguez, A. and Marcos, C.: Product user manual for the “convective rainfall rate” (CRR – PGE05 v3.1.1), NWC SAF and AEMET, available at: www.nwcsaf.org (last access: 22 March 2012), 32 pp., 2012.
- Shchepetkin, A. F. and McWilliams, J. C.: The regional ocean modelling system: a splitexplicit, free-surface, topography-following-coordinate oceanic model, *Ocean Model.*, 9, 347–404, 2005.
- Signell, R. P., Carniel, S., Cavaleri, L., Chiggiato, J., Doyle, J. D., Pullen, J., and Sclavo, M.: Assessment of wind quality for oceanographic modelling in semi-enclosed seas, *J. Marine Syst.*, 53, 217–233, 2005.
- Smolarkiewicz, P. K. and Margolin, L. G.: MPDATA: A finite-difference solver for geophysical flows, *J. Comput. Phys.*, 140, 459–480, 1998.
- Stanešić, A.: Assimilation system at DHMZ: development and first verification results, *Cro. Met. Jour.*, 44/45, 3–17, 2011.
- Termonia, P.: Scale-selective digital filter initialization, *Mon. Weather Rev.*, 136, 5246–5255, 2008.
- Thoss, A.: Product user manual for SAFNWC/MSG “precipitating cloud” (PC – PGE04 v1.5), NWC SAF and SMHI, available at: www.nwcsaf.org (last access: 22 March 2012), 18 pp., 2012.
- Torres, O., Decae, R., Veefkind, J. P., and de Leeuw, G.: OMI Aerosol Retrieval Algorithm, Clouds, Aerosols and Surface UV Irradiance, edited by: Stammes, P., vol. III, Version 2.0, 2002.
- Tudor, M. and Ivatek-Šahdan, S.: The case study of bura of 1 and 3 February 2007, *Meteorol. Z.*, 19, 453–466, 2010.
- Tudor, M. and Termonia, P.: Alternative formulations for incorporating lateral boundary data into limited area models, *Mon. Weather Rev.*, 138, 2867–2882, 2010.
- Veihelmann, B., Levelt, P. F., Stammes, P., and Veefkind, J. P.: Simulation study of the aerosol information content in OMI spectral reflectance measurements, *Atmos. Chem. Phys.*, 7, 3115–3127, doi:10.5194/acp-7-3115-2007, 2007.

- Wolf, J. and Luksch, J.: Bericht an die Seebehörde in Fiume über die Vorexpedition Nautilus im Sommer 1874 zum Zwecke physikalischer Untersuchungen des adriatischen Meeres, I, Bericht, Fiume, 55 pp., 1887.
- Zaninović, K., Gajić-Čapka, M., Perčec Tadić, M., Vučetić, M., Milković, J., Bajić, A., Cindrić, K., Cvitan, L., Katušin, Z., Kaučić, D., Likso, T., Lončar, E., Lončar, Ž., Mihajlović, D., Pandžić, K., Patarčić, M., Srnec, L., Vučetić, V.: Climate Atlas of Croatia: 1961–1990: 1971–2000, CMHS Monograph, Zagreb, 200 pp., 2008.

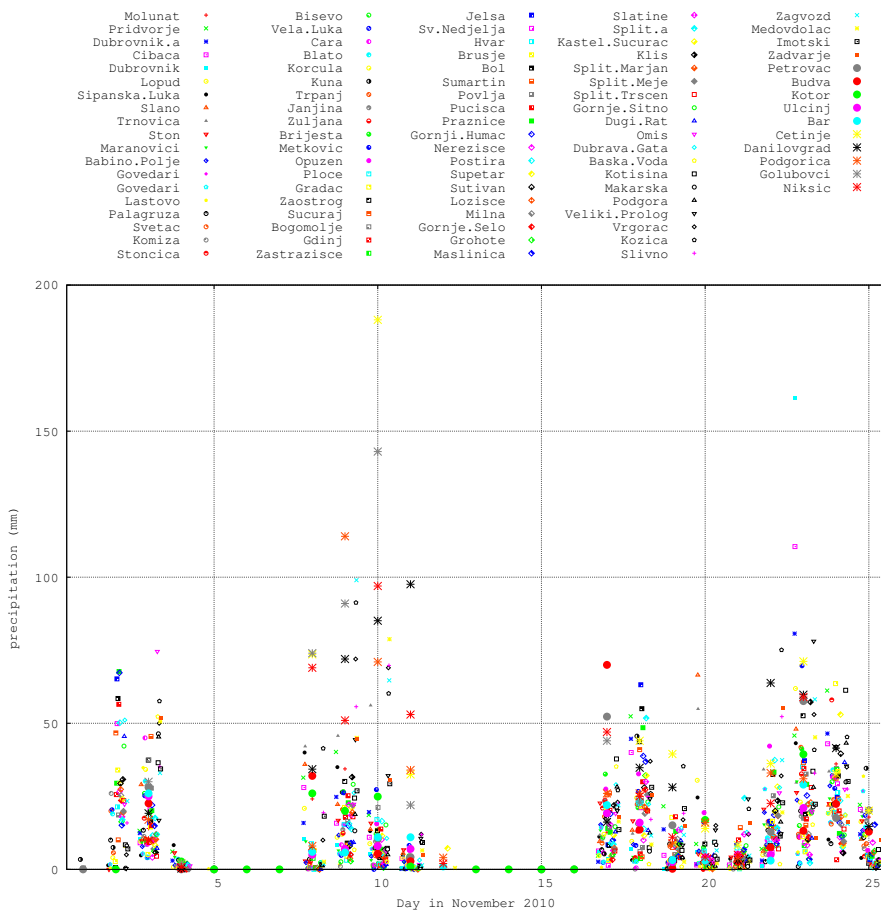


Figure 2. Measured accumulated 24 hourly precipitation on rain ~~guages~~-gauges in southeastern Croatia (smaller ~~sybol~~symbols) and ~~Montenegroi~~-Montenegro (larger symbols) during November 2010. Precipitation shown for a certain day is measured at 6 UTC accumulated from the previous 24 hours. Only stations close to Adriatic coast are shown.

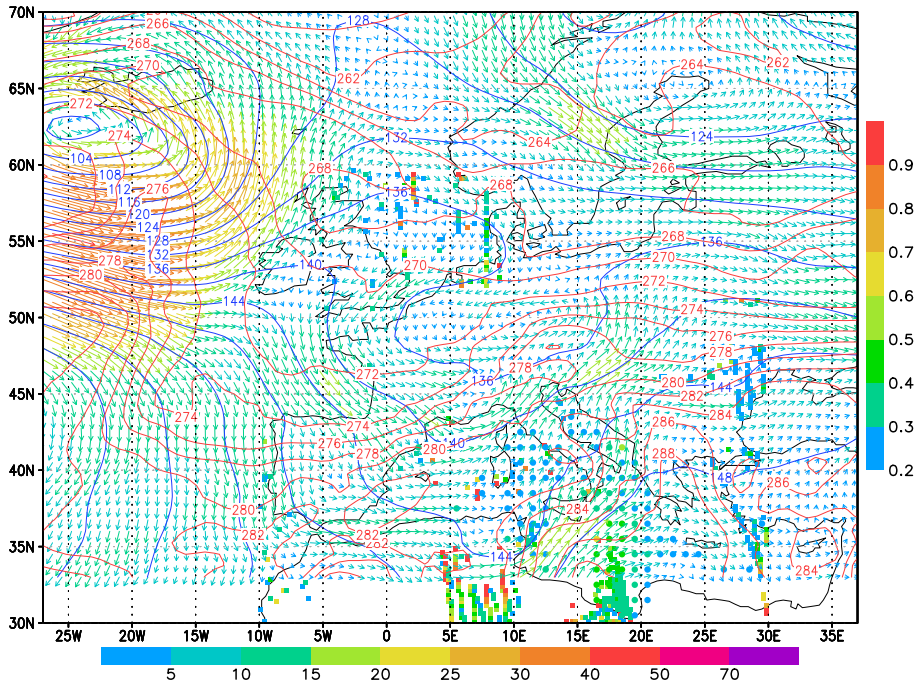


Figure 3. ERA Interim 850 hPa wind (colour of the vectors shows wind speed in m/s as on the colour bar below), geopotential (blue isolines) and temperature (red isolines) with measured aerosol optical thickness at 12:00 UTC 7 November 2010 from MODIS (circles) and OMI (squares). AOT is shown in colour according to the colour bar on the right.

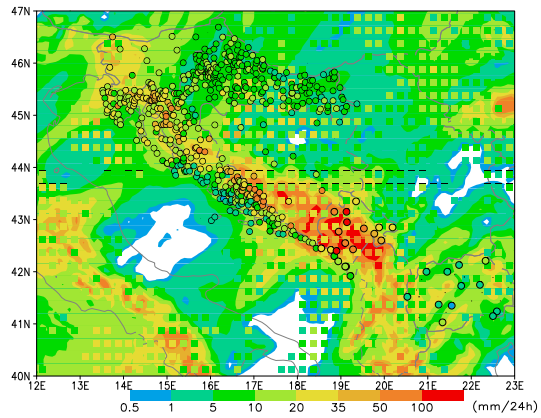
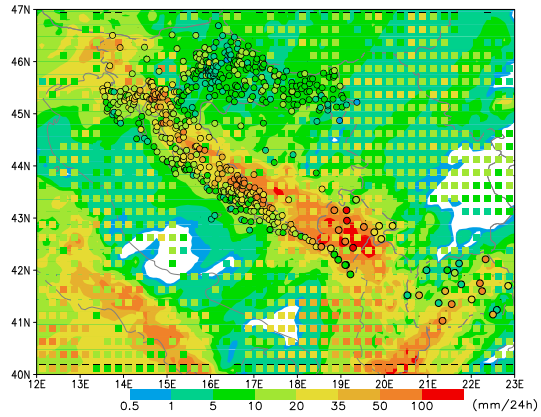


Figure 4. Measured accumulated 24 hourly precipitation on rain [gauges-gauges](#) in Croatia, Montenegro and Macedonia (circles), TRMM rainfall data (squares) and 8 km ALADIN forecast data (shaded [background](#)), [the precipitation is accumulated](#) for the period from 06:00 UTC on 8 [until 06 UTC on 9 \(top\)](#) and [from 06 UTC on 9 until 06 UTC on 10 \(bottom\)](#) November 2010. [The corresponding rain-gauge data is also shown in Fig. 2 as measured on 9 and 10 November 2010.](#)

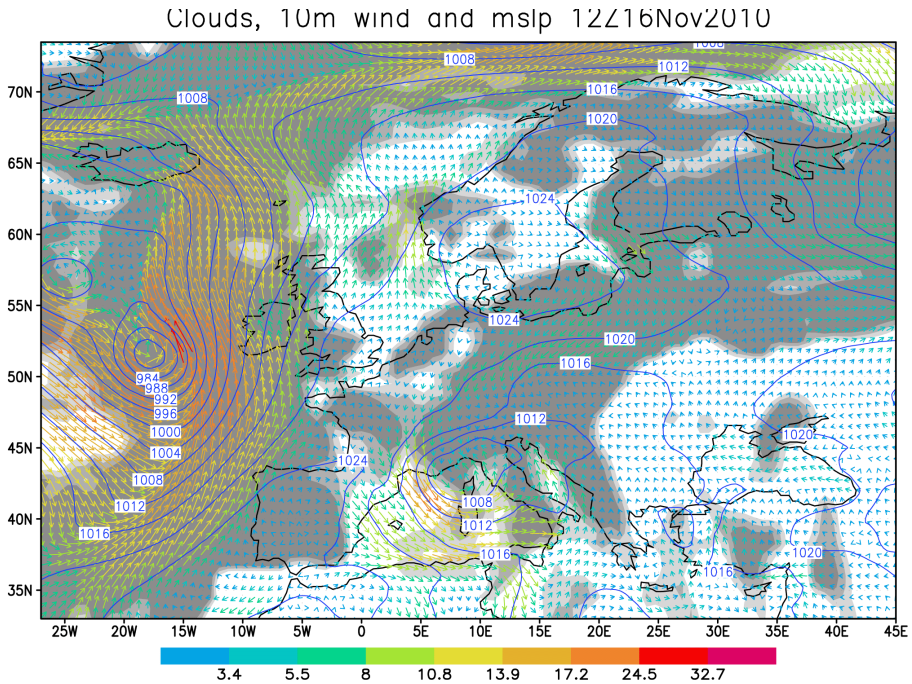


Figure 5. ERA Interim 10 m wind ([colour of the vectors shows wind speed in m/s as on the colour bar below](#)), mean sea level pressure (blue [isolines](#)) and cloudiness (shades of gray) [for 12 UTC on 16 November 2010](#).

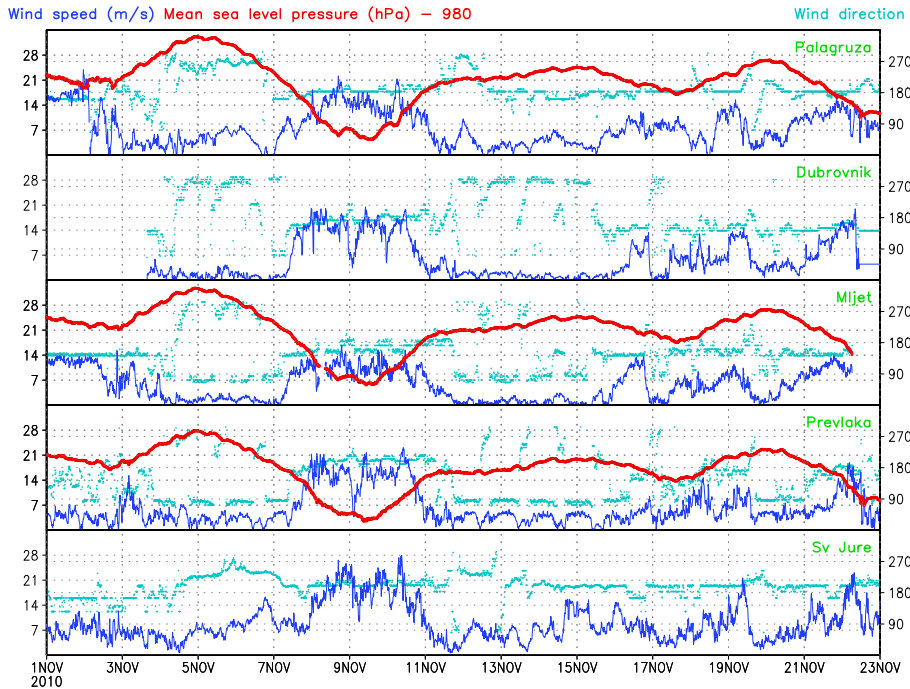


Figure 6. Measured wind speed (dark blue, [scale on the left](#)) and mean sea level pressure (red) reduced by 980 hPa ([to fit the scale on the left](#)) and wind direction (light blue, [in degrees from north clockwise, scale on the right](#)) for November 2010. [Sv. Jure is marked as Biokovo in Fig. 1.](#)

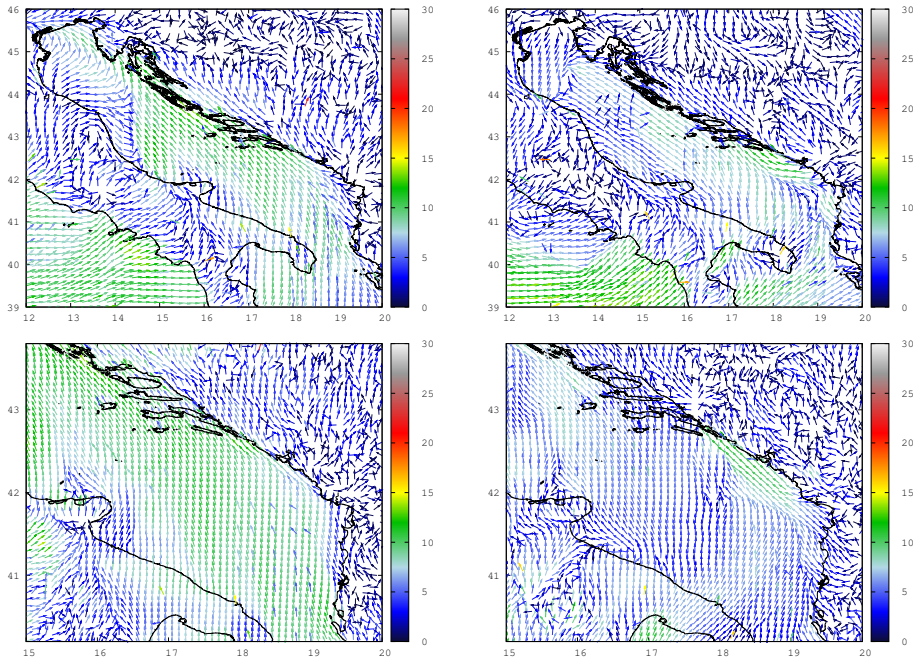


Figure 7. Forecast 10 m wind in 8 (top) and 2 km (bottom) resolution and measured wind speed and direction (arrows) from MetOp ASCAT data (above the sea surface), SYNOP and automatic stations for 12:00 UTC 16 (topleft) and 17 (bottomright) November 2010. Colour of the vectors shows wind speed in m/s as on the colour bars. Model data are shown as thin vectors on a denser grid in the background, measured wind is shown as thicker vectors on the location of measurement.

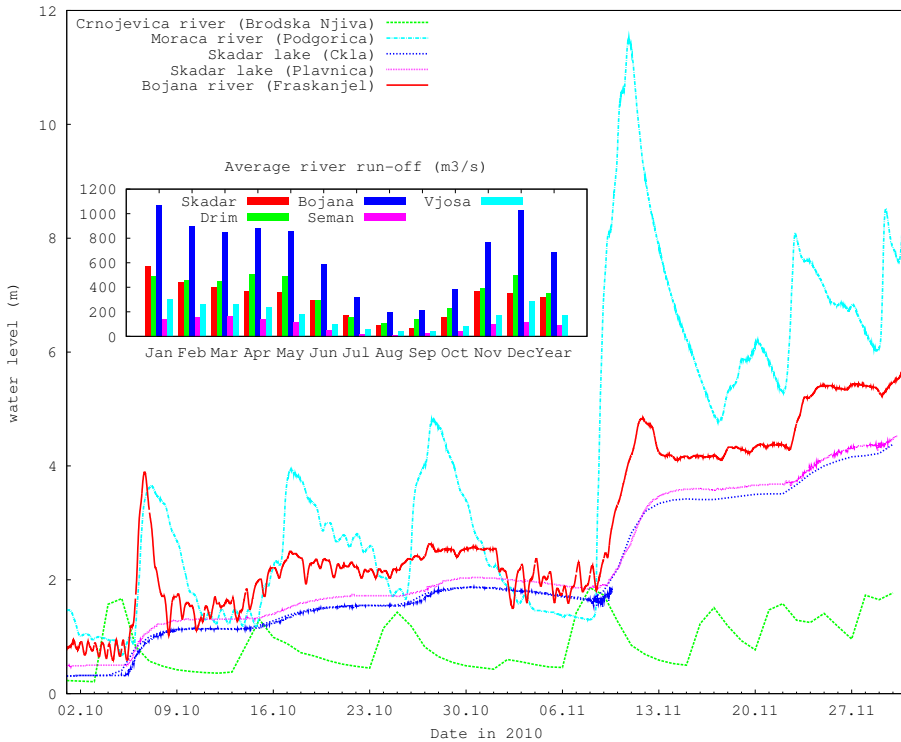


Figure 8. Measured water levels of rivers in Montenegro and Skadar lake in October and November 2010 and climatological river run-off of Albanian rivers (embedded figure). [Moraca and Crnojevica rivers are tributaries to the Skadar lake, Bojana river takes the outflow from the Skadar lake and flows into the Adriatic sea at the position marked as Bojana in Fig 1.](#)

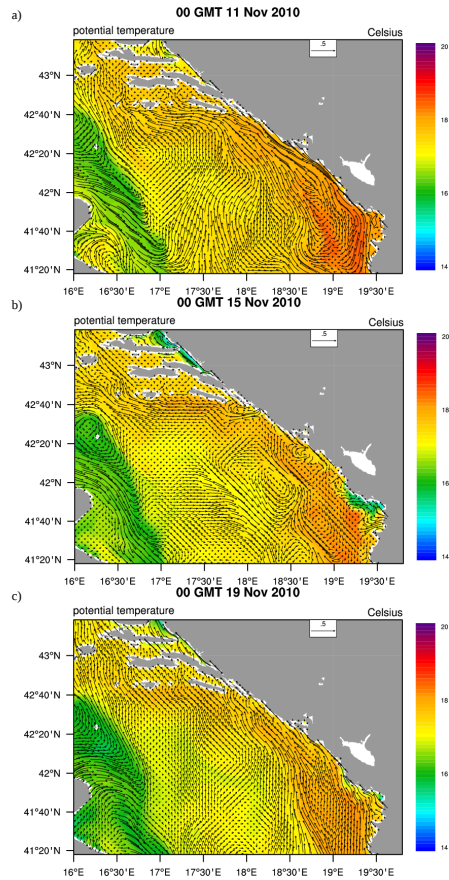


Figure 9. Surface currents (vectors) and sea surface temperature (shaded background) from ROMS for 00:00 UTC on 11 (a), 15 (b) and 19 (c) November 2010.

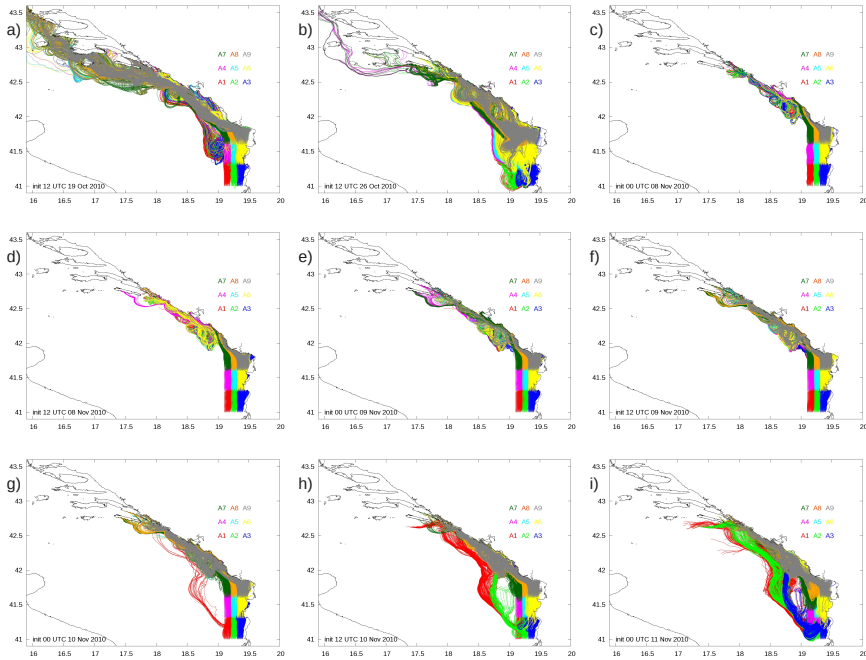


Figure 10. Trajectories of drifters released at 12:00 UTC 19 (a) and 12:00 UTC 26 October (b), 00:00 (c) and 12:00 (d) UTC 8, 00:00 (e) and 12:00 (f) UTC 9, 00:00 (g) and 12:00 (h) UTC 10 and 00:00 UTC 11 (i) November 2010. [Trajectories initiated off different parts of a rectangle in southeast Adriatic are plotted in different colours.](#)